Wind Tunnel Measurements of Pollutant Plume Dispersion over Hypothetical Urban Areas

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Revised manuscript BAE-D-17-02025R1 submitted to

Building and Environment on January 31, 2018

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Abstract

Non-computational-fluid-dynamics (Non-CFD) solutions, such as Gaussian plume models, 10 are commonly employed to predict ground-level pollutant concentrations because of their cost-11 effectiveness. Whilst, they should be applied with caution for pollutant plume dispersion over 12 complicated urban morphology in view of their implicit limitation of empirically determined 13 14 dispersion coefficients σ_z . Skin-friction coefficient c_f , which is a measure of aerodynamic resistance induced by rough surfaces, is proposed to parameterize the dispersion coefficient over 15 urban areas in isothermal conditions. Analytical derivation shows that σ_z is proportional to the 16 newly proposed friction length scale $L_f(=x^{1/2} \delta^{1/2} c_f^{1/4}$ where x and δ are the distance after pollutant 17 source and the turbulent boundary layer thickness, respectively). Its functional form is verified by 18 wind tunnel experiments for flows and tracer plume dispersion over hypothetical urban areas in 19 the form of idealized street canyons of different building-height-to-street-width (aspect) ratios 20 (ARs = 1/2, 1/4, 1/8 and 1/12). A ground-level, pollutant line source in crossflows is modeled by 21 atomizing water vapor using ultrasonic. Ranges of turbulent boundary layer thickness (240 mm ≤ 22 $\delta \leq 285$ mm) and skin-friction coefficient (8×10⁻³ $\leq c_f \leq 13\times10^{-3}$) are tested. The tracer 23 concentrations over rough surfaces exhibit the Gaussian distribution. A close correlation between 24 σ_z and L_f is revealed (coefficient of determination $R^2 = 0.93$), demonstrating the influence of drag 25 on the transport processes. The analytical solution and wind tunnel results collectively suggest an 26 improved parameterization of pollutant plume dispersion coefficient over rough surfaces, refining 27 28 the practice of the air quality forecast in urban areas. (Word count: 251)

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Keywords: Aerodynamic resistance; hypothetical urban areas; pollutant plume dispersion; rough
 surfaces; vertical dispersion coefficient.

32 **1. Introduction**

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Rapid urbanization and expanding human activities are accompanied with increasing pollutant emission and degrading air quality [1]. Roads in cities are flanked by closely packed, high-rise buildings, forming arrays of street canyons. The dynamics in the urban canopy layer (UCL) are different from those in the atmospheric boundary layer (ABL) [2]. Moreover, the drag induced by ground surfaces modifies the wind flows and the pollutant transport aloft [3,4]. Advanced understanding of the pollutant transport processes over urban areas is utmost important for public health and the formulation of pollution control strategy.

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Gaussian plume model is the well-received solution to ABL pollutant dispersion because 42 of its cost-effectiveness in air quality management [5]. Its accuracy, however, tightly depends on 43 the functionality of dispersion coefficients (σ_y in lateral and σ_z in vertical direction). Theoretically, 44 dispersion coefficients are functions of atmospheric turbulence and surface roughness [6,7]. Most 45 σ_{v} and σ_{z} are empirically determined based on the atmospheric stability and the distance behind the 46 pollutant source [8]. Whereas, the turbulence in the atmospheric surface layer (ASL) is 47 complicated by the land feature such as natural terrain, forest vegetation or building geometry. It 48 in turn modifies substantially the dispersion coefficients (especially σ_z), which, however, is often 49 50 overlooked in the practice of pollutant plume dispersion modeling [9].

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Numerous studies have been conducted to enrich our understanding of tracer transport over 52 rough surfaces [10,11]. Reduced-scale physical modeling, such as using wind tunnels [12,13] or 53 water channels [14,15], was one of the commonly adopted approaches. Various influential factors 54 on pollutant dispersion, such as thermal effect [16], upstream and downstream buildings [17-19], 55 roof-top structures [20,21], roof pitches [22] and roadway configurations (noise barriers, roadway 56 elevation and nearby terrain) [23], were examined by wind tunnel experiments. Plume dispersion 57 over rough surfaces was also widely investigated. Salizzoni et al. [24] compared the plume 58 dispersion behavior over an array of (identical) roughness elements of different density together 59 with the sensitivity to the existence of smaller roughness scales. It was found that the turbulent 60 mass flux increased with decreasing obstacle density. Moreover, those small roughness scales 61 62 further enhanced the turbulent transport processes only in the skimming flow regime (closely

packed roughness elements). To study the dense-gas dispersion over rough surfaces, a series of 63 wind tunnel experiments were conducted in three different wind tunnels in a cooperative program 64 [25]. By monitoring the vertical concentration profiles of dense gas (carbon dioxide CO₂) and non-65 buoyant tracer (ethane C₂H₆) at different locations after line sources in crossflows, detailed data 66 were archived to refine the vertical-turbulent-entrainment component of dispersion models [26,27]. 67 Recently, Perry et al. [28] tested the tracer dispersion over elongated rectangular buildings. The 68 vertical and lateral profiles of concentrations were measured and the Gaussian distribution was 69 applied well in the regression. It was found that the effective plume height descended faster when 70 71 the approaching flows and the buildings were inclined at 45°, suggesting that wind direction was 72 an important factor governing the transport processes. Apart from wind tunnel experiments, 73 mathematical modeling [29-32] and field measurements [33-35] have been deployed to investigate 74 pollutant dispersion over different configurations of roughness elements. The aforementioned studies collectively showed that pollutant plume dispersion was closely related to surface 75 morphology so diversified pollutant transport characteristics were unveiled over different 76 arrangements of obstacles. Whereas, the fundamental understanding and the quantitative 77 correlation between ASL pollutant dispersion and ground-surface roughness are rather limited. 78 79

The authors have persisted to address the aforementioned scientific questions for years. 80 81 Water evaporation was employed to compare the capacity of ventilation and pollutant removal from street canyons of different building-height-to-street-width (aspect) ratios (ARs) [36]. Skin-82 friction coefficient c_{f} , as a measure of urban roughness, has been adopted for the parameterization 83 of ventilation and pollutant removal using both numerical modelling [7,37] and wind tunnel 84 measurements [38-40]. As an extension of our ongoing research effort, this paper looks into the 85 flows and transport processes over hypothetical urban areas. We attempt to parameterize the 86 vertical dispersion coefficient σ_z in the (conventional) Gaussian plume model by formulating the 87 functional form in terms of skin-friction coefficient c_f and other flow variables. We suggest the 88 use of Gaussian-plume framework for the concentration profiles because of its solid theoretical 89 basis. Our hypothesis is to fine-tune the dispersion coefficient to account for the effect of urban 90 morphology on transport processes. The theory is derived first in the next section. Afterward, wind 91 92 tunnel results are reported to verify the mathematical hypothesis and to characterize the tracer 93 plume dispersion as a function of surface roughness. This study advances our understanding of 94 pollutant transport mechanism over cities in response to urban morphology as well as
95 parameterizes the ASL pollutant-plume (vertical) dispersion coefficient over urban areas.

96

97 2. Theoretical background

- 98 2.1. Skin-friction coefficient
- 99

100 Skin-friction coefficient

$$c_{f} = \frac{\tau_{w}}{\rho U_{\infty}^{2}/2} = \frac{2u_{*}^{2}}{U_{\infty}^{2}},$$
(1)

101 which is defined as the ratio of drag to (half of) dynamic pressure, is commonly used to measure the aerodynamic resistance for flows over (non-smooth) solid boundaries in the engineering 102 community [41,42]. Here, τ_w is the shear stress induced by the rough surface, ρ the fluid density 103 and U_{∞} the freestream wind speed [7,38]. In this paper, u_* is the friction velocity estimated by 104 extrapolating the vertical profile of turbulent momentum flux from the inertial sublayer (ISL) down 105 to the displacement height at wall-normal distance z = d [43, 44]. The ISL is defined as the region 106 where the turbulent momentum flux is rather constant (less than 10% spatial variation in this study) 107 108 [39].

109

110 *2.2. Gaussian plume model*

111

Gaussian plume dispersion model is well received by the industry to estimate ABL pollutant concentrations [8]. After a continuous, infinite line source of passive and inert pollutants in crossflows, the steady-state Gaussian plume dispersion model is

$$\psi(x,z) = \frac{Q}{\sqrt{2\pi}U\sigma_z} \left\{ \exp\left[-\frac{(z-z_s)^2}{2\sigma_z^2}\right] + \exp\left[-\frac{(z+z_s)^2}{2\sigma_z^2}\right] \right\}$$
(2)

115 where ψ is (mean) pollutant concentration, U the (uniform) wind speed, z the height measuring 116 from the ground surface, z_s the effective source height, σ_z the (vertical) dispersion coefficient and 117 Q the pollutant emission rate. It is noteworthy that the effective source height z_s could be estimated 118 by the best fit of Gaussian model to the measured concentrations [4]. In this paper, the emission 119 height is essentially at the roof level so $z_s = h$ is assumed in the following analyses. It is because the flows are recirculating but not moving in the streamwise direction in-between roughness elements below the roof level. The current ground-level line source indeed would further elevate the effective source height (in the order of roughness element size *h* because of the discharge momentum). The related uncertainty is discussed in *Section 3.2 Source design*. The vertical profiles of pollutant concentration at any downstream positions *x* after the pollutant source are in the forms of Gaussian distribution characterized by the dispersion coefficient σ_z . Therefore, the reliability of σ_z is crucial to the accuracy of Gaussian plume dispersion models.

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128 2.3. Aerodynamic resistance and plume dispersion

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130 Dispersion coefficient σ_z is a function of ABL turbulence, surface roughness and the 131 distance after the pollutant source [6]. It can be described by the classic *K*-theory, anti-gradient 132 diffusion model, as follows

$$\sigma_z^2 = 2K_z t = 2K_z \frac{x}{U}$$
(3)

where K_z is the eddy diffusivity in the vertical direction *z* and *t* (= x/U) the pollutant traveling time from the source to the receptor in the streamwise direction [45]. In the context of mixing-length theory, the eddy diffusivity can be approximated by

$$K_z \propto l_* u_* \tag{4}$$

136 where l_* and u_* (friction velocity) are the characteristic scales of length and velocity, respectively, 137 of turbulent eddies. For the ground-level source, the pollutant transport is mainly governed by the 138 near-surface turbulence so the friction velocity u_* is used as the velocity scale for the eddy 139 diffusivity. Without loss of generality, the length scale l_* is comparable to the turbulent boundary 140 layer thickness δ . Equation (3) is then simplified to

$$\sigma_z^2 \propto 2x\delta \frac{u_*}{U} \,. \tag{5}$$

141 Substituting Equation (1) $u^{*}/U_{\infty} = (c_{f}/2)^{1/2}$ into Equation (5) yields the basic functional form

$$\sigma_z \propto x^{1/2} \times \delta^{1/2} \times c_f^{-1/4}.$$
 (6)

143 Alternatively, Equation (6) can be derived from the Lagrangian approach. The (streamwise) 144 distance traveled by advection in time t is $x (= U \times t)$. By first-order approximation, the 145 corresponding (vertical) distance traveled by diffusion is $z = (K_z \times t)^{1/2}$. The distance travel in the 146 streamwise x and vertical z directions is therefore correlated by

$$z^2 = K_z \frac{x}{U}.$$
 (7)

We apply the skin-friction coefficient Equation (1) and the mixing-length theory Equation (4) 147 again, Equation (7) is simplified to $z^2 = x l^* c f^{1/2} / 2^{1/2}$. Note that the eddy size is comparable to the 148 turbulent boundary layer thickness $l^* \approx \delta$ and the vertical dispersion coefficient is proportional to 149 the vertical distance traveled by pollutant $\sigma_z \propto z$, Equation (6) is thus arrived as well. Calder [46] 150 proposed the analytical solution that $K_z = \kappa u \cdot z$. Hence, σ_z can be calculated as $\sigma_z^2 = 2 K_z t = 2 \kappa$ 151 $u \ge z x/U$ based on Equation (3) and the height z is introduced into the calculation of dispersion 152 coefficient σ_z . However, the functional form is different from that of conventional Gaussian models. 153 Arya [45] used the drag coefficient to estimate K_z but the characteristic length scale l^* was not 154 considered. Britter et al. [47] used the entrainment velocity w_e to replace u^* ($w_e = 0.65u^*$ based on 155 wind tunnel results) while σ_z and w_e is proportional to the rate of change of vertical dispersion 156 coefficient $w_e \propto d\sigma_z/dt$. We also tested using the size of roughness elements h to estimate the length 157 scale but unsuccessful. Afterward. We hypothesized that the transport processes are dominated by 158 the large-scale eddies so use the boundary layer thickness δ as the length scale, arriving the current 159 formulation. However, our understanding of the relation between turbulence length scale l_* and 160 the (parameterization of) dispersion coefficient σ_z is still limited. Additional studies are being 161 undertaken to address the questions. In this paper, the analogous skin-friction coefficient and the 162 boundary layer thickness is employed to parametrize σ_z in order to keep the conventional 163 formulation of dispersion coefficient in terms of streamwise distance x only but not vertical 164 distance z. It is thus proposed that dispersion coefficient can be measured by the aerodynamic 165 166 resistance induced by roughness elements, which, however was not verified in the previous studies.

167

168 **3. Methodology**

169 *3.1. Wind tunnel setup*

Laboratory-scale experiments are performed in the open-circuit wind tunnel in the 171 Department of Mechanical Engineering, The University of Hong Kong (Fig. 1). A honeycomb-172 type filter is installed upstream to suppress the (background) turbulence level due to the U-shape 173 duct before the wind-tunnel test section. The wind-tunnel test section is 6-m long, 0.56-m wide 174 and 0.56-m high. The design wind speed is in the range of 0.5 m sec⁻¹ $\leq U \leq 15$ m sec⁻¹. To model 175 a turbulent boundary layer in the test section, a 2-m long upstream section is adopted on which an 176 array of square aluminum tubes (size h = 19 mm with separation 19 mm apart) is glued (Fig. 1a). 177 The freestream wind speed U_{∞} , which is measured by a Prandtl-type pitot-static tube installed 178 upstream of the test section, is maintained at 3.3 m sec⁻¹ and 6.6 m sec⁻¹, to examine the flow 179 independence and scale similarity. The Reynolds number Re_{∞} (= $U_{\infty} \delta / v$) is calculated based on 180 the freestream wind speed U_{∞} and the turbulent boundary layer thickness δ that is in the range of 181 $80,000 \le Re_{\infty} \le 200,000$. It is sufficiently high for negligible molecular viscosity effect. The wind 182 tunnel is equipped with a traverse system for sensor positioning which is controlled by the National 183 Instruments (NI) motion-control unit (1-mm spatial resolution). Table 1 summarizes the 184 measurement setting employed in this study. 185

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187 *3.2. Source design*

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A ground-level line source of pollutants in crossflows is positioned in the middle of one of 189 190 the street canyons. Water vapor, which is generated by an atomizer, is released as a tracer from the line source (Fig. 1b). It is driven to the ground-level line source (of width 5 mm) from a reservoir 191 below the wind tunnel test section by an electric axial fan of constant wind speed. The amount of 192 water vapor-moisture mixture is controlled by the power to the atomizer which is carefully 193 adjusted to minimize the water-moisture emission in the experiments. The near-source relative 194 humidity (RH) is sampled to ensure that it is not saturated. The flows at the exit of line source 195 induced by the fan is $w_0 (= 0.2 \text{ m sec}^{-1})$ which are small (3% to 6%) compared with the prevailing 196 flows U_{∞} (≈ 3.3 m sec⁻¹ or 6.6 m sec⁻¹) in the wind tunnel. The volumetric flux of air discharged 197 from the line source is also small (0.1%) compared with that of the core flows in the wind tunnel. 198 The water levels in the reservoir are the same throughout the experiments to ensure a constant 199 emission rate of water vapor per unit length $Q (= 2.4 \pm 0.6 \text{ g m}^{-1} \text{ sec}^{-1})$. In the theory of rectangular 200 jet of unit depth, the initial plume rise induced by the discharge momentum from a line source in 201

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crossflows can be estimated by the length scales of momentum that is in order of 0.02 m to 0.04 m (h to 2h).
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205 *3.3. Hypothetical urban models*

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Idealized models of urban morphology are fabricated in the form of identical street canyons 207 using arrays of rib-type roughness elements [39]. The aluminum square tubes, whose length L is 208 560 mm (L/h = 29), span across the entire wind tunnel. They are aligned normally to the prevailing 209 210 wind at separation b apart so the aspect ratio AR, which is adjustable to control the aerodynamic resistance, is equal to h/b (Fig. 1c). Four types of rough surface are considered in this paper whose 211 ARs are 1/2, 1/4, 1/8 and 1/12, covering the skimming-flow (AR = 1/2), wake interference (AR = 212 213 1/4) and isolated roughness (AR = 1/8 and AR = 1/12) regimes [48]. This definition is analogous 214 to the *d*- and *k*-type flows in engineering fluid mechanics [49].

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- 216 *3.4. Measurement techniques*
- 217

The origin (x = 0 and z = 0) is set at the line source to establish the concentration grid (Fig. 1). For each measurement test, nine vertical profiles of velocity and RH are measured at selected streamwise locations in the range of $10h \le x \le 67.5h$ downstream of the source (Fig. S1 of the Supplementary Material). Totally 72 vertical profiles are collected in the experiments and 71 points are measured along each profile. A 20-second delay in time is applied before water vapor sampling because of the time lag of the RH sensor response.

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3.4.1. Hot-wire anemometry

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Velocity measurements are probed by constant-temperature hot-wire anemometer (CTA) with a *X*-wire probe (Fig. 1d). The sensing element of the probe consists of a pair of 5- μ m (diameter) platinum-plated tungsten wires. By copper plating, the active length of the sensing element is 2 mm. The angle between two (cross) wires is 100° because a larger angle (> 90°) helps measure the elevated turbulence intensity in the near-wall region as well as reduce the errors due to inadequate yaw response in highly turbulent flows [43,49]. The analog CTA signal is digitized by a 24-bit NI

data acquisition module (NI 9239). Afterward, the digital data are acquired by a desktop computer 233 via NI CompactDAQ chassis with LabVIEW software. The sampling duration at each point is 50 234 sec and the sampling frequency is 2 kHz which are similar to those employed in previous studies 235 [39,43]. A total of 100,000 velocity data are obtained at each sampling point that is sufficient for 236 the repeatability of mean and fluctuating quantities [39]. The velocity calibration is based on the 237 universal calibration law [50]. The freestream wind speed U_{∞} data measured by CTA are also 238 compared with those by a Pitot-static tube. Velocity calibration based on 20 points in the range of 239 $0.5 \text{ m sec}^{-1} \leq U_{\infty} \leq 12 \text{ m sec}^{-1}$ is conducted in prior to the wind tunnel experiments. A good 240 correlation is observed (coefficient of determination by linear regression $R^2 = 0.999$) as well, 241 demonstrating the accuracy and reliability of our CTA measurement system. 242

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244 *3.4.2. Humidity sensor*

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The RH φ and temperature *T* are measured by Humidity and Temperature Sensor SHT75 (SENSIRION AG Switzerland). Another SHT75 sensor is placed upstream of the water vapor source to monitor the background RH φ_0 and temperature *T*₀ concurrently (Fig. 1a). The RH

$$\varphi = \frac{P}{P_s} \times 100\% \tag{8}$$

249 is then converted to the vapor concentration (specific humidity) ψ

$$\psi = \frac{m}{V} = \frac{M \times P_s}{R_u \times (273.15 + T_0)} \times (\varphi - \varphi_0).$$
(9)

Here, *m* is the mass of water vapor (g), *V* the air volume (m³), *P* the partial vapor pressure (Pa), *P*_s the saturation vapor pressure (Pa), *T*₀ the ambient temperature (°C), M (= 18.015 g mol⁻¹) the molar mass of water and R_u (= 8.314 Pa m³ K⁻¹ g⁻¹ mol⁻¹) the universal gas constant. At each sampling point, the duration for RH measurements is 50 sec at a frequency of 1 Hz. The uncertainty of RH and *T* measurements is about 2% so the error in vapor concentration measurements is less than 5%.

- 256 4. Results and discussion
- 257

Flows and vapor concentration data are measured for 50 sec at each sampling point. Vertical profiles of flows are repeated in the homogeneous streamwise direction *x*. Hence, statistic variables are defined to facilitate analyses. In the following sections, overbar $\overline{\bullet}$, angle bracket $\langle \bullet \rangle$ and prime $\bullet' (= \bullet - \langle \overline{\bullet} \rangle)$ represent the temporal average, spatial average and fluctuating component, respectively, to study the dynamics. Temporal average $\overline{\bullet}$ is the averaged property during the sampling duration. Spatial average $\langle \bullet \rangle$ is the averaged property in the homogeneous streamwise direction *x*.

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267

268 Fig. 2 depicts the vertical dimensionless profiles of mean and fluctuating velocities. Available velocity profiles from previous studies are also plotted in Fig.2 [51,52]. The velocity 269 components are normalized by the freestream wind speed U_{∞} and the vertical distance is 270 normalized by the turbulent boundary layer thickness δ . In this paper, the turbulent boundary layer 271 thickness is determined by the wall-normal distance where the spatio-temporal average velocity 272 converges to 99% of the freestream wind speed $\langle \overline{u} \rangle \Big|_{\tau=\delta} = 0.99 U_{\infty}$. It is in the range of $12h \le \delta \le$ 273 15h among the cases tested in this paper (Table 1) that is comparable to that in previous studies 274 [39,43]. The turbulent boundary layer thickness is not affected much by the freestream wind speed 275 (about 2% difference) because of the use of identical roughness elements. Moreover, it is peaked 276 at AR = 1/8 (δ = 285 mm). 277

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The average velocity profiles $\langle \overline{u} \rangle$ vary only slightly over different arrays of street canyons (Fig. 2a). This indicates that the dimensionless mean profiles are not significantly affected by different configurations of surface roughness in our wind tunnel experiments and the freestream wind speed U_{∞} is an appropriate characteristic velocity scale. In the ISL over non-smooth solid boundaries, the (theoretical) vertical variation of the average velocity $\langle \overline{u} \rangle$ in isothermal conditions is described by the logarithmic law of the wall (log-law) [53]

$$\left\langle \overline{u} \right\rangle = \frac{u_*}{\kappa} \ln \left(\frac{z - d}{z_0} \right) \tag{10}$$

where κ is the von Kármán constant (= 0.4 in this paper) and z_0 the roughness length [54]. The 285 friction velocity u^* is estimated by the turbulent momentum flux [39,43]. The roughness 286 parameters d and z_0 are determined by the regression of the average velocity data with the log-law 287 Equation (10). Table 1 tabulates the roughness parameters in which d and z_0 are in the range of 288 $0.43h \le d \le 1.25h$ and $0.009h \le z_0 \le 0.047h$, respectively. Alike the turbulent boundary layer 289 thickness, these values are comparable to those in the previous study measured over street canyons 290 of AR = 1/2 (d = 0.866h and $z_0 = 0.028h$) [51]. Higher values of d (d > h) are found in the cases of 291 AR = 1/2 (Case L1 and Case H1 in Table 1) and 1/4 at low wind speed (Case L2). This feature is 292 also observed in other studies [43,51], showing that the estimate to d could be higher than h by 293 curve fitting. 294

295

It is worthy to look into several (dimensionless) ratios in the flows at low (3.3 m sec^{-1}) and 296 297 high (6.6 m sec⁻¹) wind speeds to examine whether the dynamics are independent from Reynolds number (Table 1). The variations of u^*/U_{∞} , z_0/h and d/h are in the ranges of 1.8% to 4.5%, 14% to 298 42% and 9% to 21%, respectively. We therefore find that most cases are quite independent from 299 300 Reynolds number except the one of AR = 1/4. The variation in u^*/U_{∞} is small (4.5%) but large variations are found in z_0/h (42%) and d/h (21%) so it is hard to confirm Reynolds-number 301 independence for AR = 1/4. The reason is likely attributed to the uncertainty of determining z_0 and 302 d by the regression of measured wind speed profile to the log-law in the ISL where the turbulent 303 304 momentum flux is assumed to be rather constant (less than 10% spatial variation). Additional 305 wind-tunnel experiments (with higher wind speeds) are thus needed in the future studies. Analogous to the dimensionless ratio z_0/h , the Jensen number $Je (= h/z_0)$ is a primary scaling 306 quantity commonly employed in wind engineering. It is in the range of $20 \le Je \le 110$ whose 307 variation between the cases of low and high wind speeds is about 3% to 20%. 308

309

Noticeable differences are observed in the fluctuating wind components over different rough surfaces at different freestream wind speeds. For instance, fluctuating streamwise velocity $\langle \overline{u'u'} \rangle^{1/2}$ (Fig. 2b), fluctuating vertical velocity $\langle \overline{w'w'} \rangle^{1/2}$ (Fig. 2c) and turbulent momentum flux $\langle \overline{u'w'} \rangle$ (Fig. 2d). These findings suggest that the near-wall turbulence structure is strongly affected by the (shear or form) drag induced by the roughness elements so is the transport processes

[39]. We hypothesize the analogy of pollutant plume dispersion which is elaborated in Section 4.2. 315 Pollutant dispersion. The fluctuating components are normalized by the freestream wind speed so 316 scaling similarity is observed mainly in the outer region $z \ge 0.5\delta$. Given the same freestream wind 317 speed, the intensities of various fluxes generally increase with decreasing AR (widening street 318 width). Moreover, the profiles over the array of street canyons of AR = 1/8 overlap with those of 319 AR = 1/12, suggesting that peaked turbulent momentum flux appears when the building spacing 320 is wide enough. Whereas, the near-wall fluctuating streamwise velocity for AR = 1/4 is slightly 321 over its AR = 1/12 counterpart. Further experiments are thus needed to look into that regard. It is 322 noted that the vertical profiles of flows over the array of street canyons of AR = 1/12 are mainly 323 measured above the streets while others are measured above both streets and buildings (Fig. S1 of 324 the Supplementary Material). This sampling limitation would overestimate the near-wall turbulent 325 fluxes 10% to 20% at most [39]. 326

327

Comparing the fluctuating velocities measured at the two different freestream wind speeds 328 shows that the dimensionless profiles collapse well for $\langle \overline{u'u'} \rangle^{1/2}$ but differences are observed in 329 those of $\langle \overline{w'w'} \rangle^{1/2}$ and $\langle \overline{u'w'} \rangle$. The differences are more notable over rougher surfaces of larger 330 skin friction coefficient (wider street width) in the near-wall region. It is probably because the 331 friction-velocity-to-freestream-wind-speed ratio u^*/U_{∞} is larger in the tests with a higher 332 freestream wind speed $U_{\infty} \approx 6.6$ m sec⁻¹, leading to elevated fluctuating velocity components. The 333 difference (at most 20%) in the vertical fluctuating velocity $\langle \overline{w'w'} \rangle^{1/2}$ and the turbulent 334 momentum flux $\langle \overline{u'w'} \rangle$ among the wind-tunnel measurements at different freestream wind speeds 335 diminishes for $z \ge 0.5\delta$. The scale similarity is fulfilled only in the outer region because the 336 freestream wind speed is not the best characteristic velocity scale for the near-wall flows where 337 the dynamics are shear dominated. The scale dependence on shear using friction velocity u^* is 338 elaborated below. 339

340

Fig. 3 depicts the vertical profiles of wind velocity $\langle \overline{u} \rangle$ normalized by the friction velocity *u** that is expressed in term of the dimensionless wall-normal distance $(z - d)/z_0$ over different rough

surfaces in logarithmic scale. All the profiles, regardless of the freestream wind speed U_{∞} , clearly 343 exhibit the conventional log-law behavior, demonstrating that the ISL is well developed in our 344 wind tunnel measurements. The ISLs for flows over the array of street canyons of AR = 1/2 locate 345 much higher that are followed by their AR = 1/4, 1/8 and 1/12 counterparts (increasing 346 aerodynamic resistance). The diversified near-wall average velocities are induced by the roughness 347 sublayer (RSL) dynamics where the aerodynamic resistance is amplified, further slowing down 348 the flows. The average wind profiles over arrays of street canyons of ARs = 1/8 and 1/12 overlap 349 closely with each other because the aerodynamic resistance is dominated by the form drag (or flow 350 impingement) across individual roughness elements when their separation is wide enough 351 (prevailing flows entrain down to the street level). The difference among the cases with different 352 353 freestream wind speeds U_{∞} is small (at most 5%) because the Reynolds number Re_{∞} is sufficiently high for fully developed flows (Table 1). 354

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356 4.2. Pollutant dispersion

357

Vertical profiles of mean concentration are measured at different streamwise locations xafter the line source. As observed in the laboratory, the plume spreads gradually in the streamwise direction (Fig. S2). Fig. 4 illustrates the self-similar vertical dimensionless profiles of pollutant concentrations over arrays of street canyons of ARs = 1/2, 1/4, 1/8 and 1/12. The wall-normal distance is normalized by the dispersion coefficient

$$\sigma_{z}(x) = \begin{bmatrix} \int_{h}^{\delta} z^{2} \times \overline{\psi}(x, z) dz \\ \int_{h}^{\delta} \overline{\psi}(x, z) dz \end{bmatrix}^{1/2}$$
(11)

363 while the pollutant concentration is normalized by its canopy-level value

$$\overline{\psi}_{\text{canopy}}(x) = \overline{\psi}(x, z = h) = \sqrt{\frac{2}{\pi}} \times \frac{Q}{U\sigma_z(x)}.$$
(12)

This approach helps focus on the dynamics within plume coverage. Although it is not the maximum, the canopy-level pollutant concentration is an appropriate characteristic scale for normalization. Unlike the maximum concentration, $\overline{\psi}_{canopy}$ is independent from the elevation z.

that can be derived analytically. The vertical dimensionless profiles of concentration 367 $\overline{\psi}(x,z)/\overline{\psi}_{\text{canopy}}(x)$ measured in different streamwise positions x, regardless of freestream wind 368 speeds U_{∞} , collapse well with each other that fit into the theoretical Gaussian distribution Equation 369 (2). It is hence suggested that ASL pollutant plume dispersion over rough surfaces, such as urban 370 areas, exhibits the conventional Gaussian form similar to its ABL counterpart. A notable elevation 371 of vapor concentration is observed in the near-wall region ($z \le 0.75 \times 2^{1/2} \sigma_z$). This peaked 372 concentration is mainly attributed to the flows over the windward wall of roughness elements 373 374 where the sharp edges (singularities) enhance the local turbulence levels that in turn dilute the tracer adjacent to solid boundaries. This phenomenon was also reported by previous studies 375 [4,19,23]. The sharp mean-wind-speed gradient over rough surfaces transports pollutants by 376 advection in a non-uniform manner. Moreover, the (vertical) discharge momentum at the ground 377 level line source induces initial plume rise that ends up in elevated concentration peak in the near-378 wall region. Although the uncertainty is not large, additional tests for the sensitivity of plume 379 trajectory and tracer distribution to the source-level discharge momentum will be performed to 380 confirm the findings in this paper and the uncertainty induced by imperfect experimental setting. 381 382 Another possible reason is that the water vapor condenses on the (cooler) aluminum rib surfaces. It subsequently reduces canopy-level tracer concentrations and eventually raises the dimensionless 383 384 tracer concentrations aloft. Besides, Fig. 4a compares the vertical dimensionless profiles of tracer concentrations over the array of street canyons of AR = 1/2 between this study and Salizzoni et al. 385 [24]. The plume core converges to the theoretical Gaussian distribution. Comparing the Gaussian 386 profiles and wind tunnel measurements, the coefficient of determination R^2 is over 0.95 and the 387 root-mean-square (RMS) error is less than 0.2 so their agreement is favorable. Instead of the roof-388 level concentration $\overline{\psi}_{canony}$, MacDonald et al. [4] used the (elevated) peak concentration in the 389 normalization in which their coefficient of determination is similar to ours but the RMS error is 390 slightly improved. Nonetheless, we suggest to use the roof-level concentration and the Gaussian 391 model in the normalization in this paper because of the solid theoretical basis. Alternatives will be 392 393 tested in future studies.

Fig. 5 attempts to characterize the tracer plume dispersion over hypothetical urban areas of different aerodynamic resistance by comparing the dimensionless profiles of vertical dispersion

coefficient σ_z at different streamwise locations x, skin-friction coefficient c_f and freestream wind 397 speeds U_{∞} . The vertical dispersion coefficient σ_z increases with increasing streamwise distance x 398 because of crosswind turbulent transport together with widening plume coverage. A good fit of 399 power regression is obtained (coefficient of determination $0.980 \le R^2 \le 0.995$) for all the cases of 400 low (Fig. 5a) and high (Fig. 5b) wind speeds in which the regression power is in the range of 0.39 401 $\leq n \leq 0.47$ (Table 2). The value of *n* obtained in this paper is smaller ($\leq 25\%$) than the theoretical 402 one (n = 0.5) commonly used in the Gaussian plume models in isothermal conditions in far field 403 [5]. The smaller values of power *n* observed in the current wind tunnel measurements are mainly 404 405 attributed to the under estimate of σ_z in the far field ($x \ge 37.5h$) because of vapor condensation in the near-wall region (Fig. 5). The power *n* of σ_z regression from the LES of Wong and Liu [7] is 406 higher (0.52 $\leq n \leq$ 0.66) and of Ng and Liu [55] is lower (0.40 $\leq n \leq$ 0.58). Deviations of the 407 mathematical modeling results are probably caused by inaccurate boundary conditions and coarse 408 spatio-temporal resolution. The value of n = 0.69 over an array of street canyons of AR = 1/2409 from the wind tunnel results of Salizzoni et al. [24] is even higher. The discrepancy could be 410 attributed to the different (background) turbulence levels and/or modeling parameters in various 411 412 studies. Moreover, addition of small roughness elements at the top of the bars enhanced the turbulent fluxes that would further enlarge the dispersion coefficient. Nonetheless, the value of 413 power *n* obtained in this study is comparable to that of other studies and it generally falls into a 414 reasonable range. 415

416

The vertical dispersion coefficients σ_z measured over different rough surfaces do not 417 converge onto a universal function of streamwise distance x, demonstrating that crosswind plume 418 dispersion also depends on the surface roughness (Fig. 5). The plume spreads in the isolated 419 roughness regime (ARs = 1/8 and 1/12) and wake interference regime (AR = 1/4) are obviously 420 wider than that in the skimming-flow regime (AR = 1/2) in both U_{∞} = 3.3 m sec⁻¹ (Fig. 5a) and U_{∞} 421 $= 6.6 \text{ m sec}^{-1}$ (Fig. 5b). It is noteworthy that, within the isolated roughness regime, the plume 422 spread over the arrays of street canyons of ARs = 1/8 and 1/12 is similar. Its coverage is even 423 narrower with further increase in street width (decreasing AR) in the 3.3-m-sec⁻¹ dataset. This 424 425 feature is analogous to that of skin-friction coefficient c_f , suggesting the possible relation between the two parameters. Moreover, Salizzoni et al. [24] reported that, with increasing aerodynamic 426 resistance (from AR = 1 to 1/2), turbulent mixing is enhanced that in turn enlarges the plume 427

428 coverage. Using LES, Wong and Liu [7] showed that the vertical dispersion coefficient σ_z initially 429 increases with decreasing AR (skimming flow regime), reaches a peak at AR = 1/10 (wake 430 interference regime) then decreases thereafter with further decrease in AR (isolated roughness 431 regime). These findings are generally in line with those available in literature.

432

433 *4.3. Dependence of plume spread on aerodynamic resistance*

434

435 In most pollutant plume dispersion theory, vertical dispersion coefficient σ_z can be 436 expressed as a function of streamwise distance *x*

$$\sigma_z = Ax^n \tag{13}$$

where A and n are empirical constants. Neither A nor n consider the effects of aerodynamic 437 resistance over urban areas in the conventional Gaussian plume dispersion theory. Given the 438 different turbulence-generation mechanism in isothermal ASLs and thermally stratified ABLs, the 439 skin-friction coefficient c_f of hypothetical urban areas is introduced to handle the influence of 440 441 surface roughness and blockage orientation on the dynamics in term of σ_z . The skin-friction coefficient deduced from the extrapolation of vertical profiles of turbulent momentum flux 442 measured in different streamwise locations increases with decreasing AR, reaches the maximum 443 at AR = 1/8 and decreases thereafter for both measurements in freestream wind speeds $U_{\infty} = 3.3$ 444 m sec⁻¹ and 6.6 m sec⁻¹ (Table 1). This feature is similar to the numerical results using LES [56] 445 446 and Reynolds-averaged Navier-Stokes (RANS) turbulence models [38,57] in which a plateau of skin-friction coefficient over surface configurations close to AR = 1/8 was clearly shown. Alike 447 its skin-friction coefficient counterpart, the vertical dispersion coefficient σ_z for tracer plume 448 449 dispersion over arrays of street canyons is peaked at AR = 1/8 though the results over arrays of street canyons of ARs = 1/8 and 1/12 are very close to each other in freestream wind speed U_{∞} = 450 6.6 m sec⁻¹. This is mainly because the turbulent velocities are similar between these two surface 451 configurations (Fig. 1). 452

453

454 Section 2. Theoretical background derives analytically the vertical dispersion coefficient 455 σ_z as a function in terms of the downwind distance measured from the pollutant source *x*, the 456 turbulent boundary layer thickness δ and the skin-friction coefficient c_f i.e. Equation (6) $\sigma_z \propto$ 457 $x^{1/2} \times \delta^{1/2} \times c_f^{1/4}$. Most existing studies available in literature used n = 0.5. To verify the analytical

hypothesis, the vertical dispersion coefficient σ_z is plotted against the newly proposed friction 458 length scale L_f (= $x^{1/2} \times \delta^{1/2} \times c_f^{1/4}$; Fig. 6). The length dimensions (σ_z , x and δ) are normalized by the 459 size of roughness elements h. The vertical dispersion coefficient is almost directly proportional to 460 the friction length scale $\sigma_z \propto L_f$. The nonzero $\sigma_z \approx 0.4 h$ at $L_f = 0$ is caused by the lifting of 461 effective source height and the x offset due to the physical width of the line source. Scale similarity 462 is clearly demonstrated even the data are collected in different freestream wind speeds. Linear 463 regression shows that there is a close relation between σ_z and L_f (coefficient of determination by 464 liner regression $R^2 = 0.93$). This analytical formulation thus could be used in the future 465 parameterization of pollutant plume dispersion over urban areas to handle complicated building 466 roughness in isothermal conditions. 467

468

469 **5.** Conclusion

470

This paper first reports the use of water vapor to study the pollutant plume dispersion over hypothetical urban areas in laboratory-scale wind tunnel experiments. The tracer concentrations over arrays of street canyons exhibit the conventional Gaussian distribution that are compared favorably with those in previous studies. It is hence suggested the feasibility of using water vapor as an affordable and harmless tracer in wind tunnel experiments. The analytical derivation and empirical solution collectively suggest an improved parameterization of pollutant plume dispersion coefficient over rough surfaces using skin friction coefficient ($8 \times 10^{-3} \le c_f \le 13 \times 10^{-3}$).

478

The (maximum) difference in the concentrations between wind-tunnel experiments and 479 Gaussian model is up to 10% to 20% that is mainly attributed to the non-uniform mean wind speed 480 and discharge momentum from the ground-level line source, which, however, do not exist in 481 idealized configuration. Although the discrepancy is not negligible, the simplicity, user-482 friendliness and cost-effectiveness of the Gaussian-model framework are attractive benefits from 483 the end-users' perspective solving practical engineering problems. The current parameterization 484 might case up to 20% underestimation of the peaked pollutant concentration. Nonetheless, studies 485 of rectification, such as including the non-zero plume trajectory, are currently undertaken to 486 improve the accuracy of the parameterization. 487

We note the rather narrow range of boundary layer thickness $(12h \le \delta \le 15h)$ tested in this 489 paper. It is attributed to the use of identical roughness elements in the wind tunnel experiments. 490 As a preliminary study, more conservative roughness elements are employed in order to maintain 491 the homogeneity in the horizontal extent. After earned a basic understanding, more diversified 492 roughness-element types, source configurations and flows conditions should be performed to 493 further verify the newly developed parameterization of (vertical) dispersion coefficient. The 494 current wind tunnel experiments are conducted in the isothermal conditions. In view of the 495 different turbulence generation mechanism, the newly developed parameterization must be applied 496 cautiously in case stable or unstable atmospheric stratification is considered. 497

498

Only 2D ribs in crossflows are adopted in this paper, which, however, are seldom found in realistic urban areas. The lack of representation would reduce the functionality of our work. Besides, prevailing wind direction normal to the street axes is only one of the many scenarios. Deviation from the normal wind would change the flows substantially that might weaken the conclusions arrived in this paper. More roughness elements of different shape and geometry are therefore needed to verify our newly developed parameterization.

505

Apart from the aforementioned limitation, future studies will focus on the validation of the analytical and empirical solutions in a systematic manner. The major benefit of the newly developed parameterization, i.e. use the simple skin-friction coefficient c_f instead of the complicated urban morphology in terms of floor plan, obstacle blockage and building orientation, etc., to model the effect of rough urban surfaces on pollutant plume dispersion by noncomputational-fluid-dynamics (non-CFD) air quality impact assessment.

512

513 Acknowledgements

514

We would like to thank the anonymous reviewers for their invaluable suggestion for the improvement of the manuscript. This study is supported by the General Research Fund (GRF) 17205314 of the Hong Kong Research Grants Council (RGC). The technical support from Mr. Vincent K.W. Lo as well as the setup of wind tunnel infrastructure by Dr. Yat-Kiu Ho are appreciated.

Mo and Liu (2018)

520	Manuscript word count: 6,800						
521	(Including text, figure captions and table legends, but not references).						
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698	are the corresponding quantities in $U_{\infty} = 6.6$ m sec ⁻¹ . Yellow-filled symbols are the wind
699	tunnel results of Salizzoni et al. [23] at $x = 9h$ (\Box); 15h (\Box). 22.5h (\Box) and 30h (\Box) in

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Supplementary Material
Fig. S1 Measurement locations in streamwise direction.
Fig. S2. Vertical dimensionless profiles of pollutant concentrations $\overline{\psi}/\overline{\psi}_{canopy}$ plotted against wall-
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= 3.3 m sec ⁻¹ . Filled symbols are the corresponding quantities in U_{∞} = 6.6 m sec ⁻¹ .

Measurement cases		Low Wind Speed ($U_{\infty} \approx 3.3 \text{ m sec}^{-1}$)				High Wind Speed ($U_{\infty} \approx 6.6 \text{ m sec}^{-1}$)			
		Case L1	Case L2	Case L3	Case L4	Case H1	Case H2	Case H3	Case H4
Aluminum rib [mm]	Size <i>h</i>	19	19	19	19	19	19	19	19
	Separation <i>b</i>	38	76	152	228	38	76	152	228
Aspect ratio $AR (= h/b)$		1/2	1/4	1/8	1/12	1/2	1/4	1/8	1/12
Boundary layer thickness δ [mm]		240	260	285	265	245	265	285	260
Wind speed [m sec ⁻¹]	Freestream U_{∞}	3.3	3.3	3.3	3.3	6.7	6.6	6.7	6.6
Friction velocity [m sec ⁻¹]	\mathcal{U}^*	0.18	0.22	0.22	0.22	0.38	0.45	0.47	0.47
	$u*/U_{\infty}$	0.056	0.065	0.068	0.068	0.057	0.068	0.071	0.071
Skin-friction coefficient	Minimum	7.3	10	12	11	7.8	11	11	12
[×10 ⁻³]	Mean	8.0	10	12	12	8.4	11	13	13
$c_f (= 2u^{*2} / U_{\infty}^2)$	Maximum	8.5	11	13	13	8.8	12	14	14
Reynolds number	$Re_{\infty} (= U_{\infty}\delta/v)$	79	86	93	84	160	180	190	170
[×10 ³]	$Re*(=u*\delta/v)$	4	6	6	6	9	12	14	12
Displacement height	d/h	1.13	1.04	0.48	0.81	1.25	0.84	0.43	0.74
Roughness length	z_0/h	0.009	0.017	0.041	0.032	0.011	0.026	0.047	0.038
Jensen number	$Je (= h/z_0)$	110	59	24	31	91	38	21	26

Table 1. Configuration of the models of hypothetical urban areas and the flows in the wind tunnel experiments.

Remark: L and H denote the cases in low wind speed ($U_{\infty} \approx 3.3 \text{ m sec}^{-1}$) and high wind speed ($U_{\infty} \approx 6.6 \text{ m sec}^{-1}$), respectively. The skin-friction coefficient c_f is rounded to two significant figures.

				Curren	t study
Aspect ratio	Salizzoni et al. [24]	Wong and Liu [7]	Ng and Liu [55]	Low wind speed	High wind speed
				$(U_{\infty} \approx 3.3 \text{ m sec}^{-1})$	$(U_{\infty} \approx 6.6 \text{ m sec}^{-1})$
1/2	0.69	0.66	0.40	0.42	0.47
1/4		0.52	0.43	0.41	0.39
1/8		0.54	0.48	0.41	0.41
1/12		0.56	0.58	0.45	0.47

Table 2. Power n of the power-law fitting measured in different settings.

Remark: Power-law fitting $\sigma_z = Ax^n$.



(b) H₂O atomizer (c) Rib configuration (d) Sensor location (e) Source location

Fig. 1. Schematic of wind tunnel setup, rough surface configuration and source design. For visualization purposes, moisture emission is captured in Fig. 1 (e). In the experiments, only water vapor is emitted from the ground-level line source by the water atomizer.



Fig. 2. Vertical dimensionless profiles of flow variables plotted against wall-normal distance (z-h)/δ over arrays of street canyons of aspect ratios ARs = 1/2 (□); 1/4 (△); 1/8 (◊) and 1/12 (◊) in freestream wind speed U_∞ = 3.3 m sec⁻¹. Filled symbols are the corresponding quantities in U_∞ = 6.6 m sec⁻¹. Also shown are the wind tunnel data available from literature: AR = 1/2 roof top (▼) [51], AR = 1/2 cavity top (▼) [51] and AR = 1/3 (▼) [52].



Fig. 3. Vertical dimensionless profiles of mean velocity $\langle \overline{u} \rangle / u_*$ plotted against wall-normal distance $(z - d)/z_0$ over arrays of street canyons of aspect ratios ARs = 1/2 (\square); 1/4 (\triangle); 1/8 (\diamondsuit) and 1/12 (\bigcirc) in freestream wind speed $U_{\infty} = 3.3$ m sec⁻¹. Filled symbols are the corresponding quantities in $U_{\infty} = 6.6$ m sec⁻¹. Also shown is the conventional logarithmic law of the wall (log-law) Equation (10) (dashed line).



Fig. 4. Vertical dimensionless profiles of tracer concentrations w/w canopy plotted against wall-normal distance (z-h)/2^{1/2}σ_z over arrays of street canyons of aspect ratios ARs = (a) 1/2; (b) 1/4; (c) 1/8 and (d) 1/12 at x = 10h (□); 15h (△); 22.5h(▽); 30h (▷); 37.5h (<); 45h (◇); 52.5h (+); 60h (-) and 67.5h (○) in freestream wind speed U_∞ = 3.3 m sec⁻¹. Filled symbols are the corresponding quantities in U_∞ = 6.6 m sec⁻¹. Yellow-filled symbols are the wind tunnel results of Salizzoni et al. [23] at x = 9h (□); 15h (□). 22.5h (□) and 30h (□) in freestream wind speed U_∞ = 6.6 m sec⁻¹. Also shown is the theoretical Gaussian-form tracer concentrations (dark solid line).



Fig. 5. Dimensionless vertical dispersion coefficients σ_z/h in the streamwise direction x/h over arrays of street canyons of aspect ratios ARs = 1/2 (\Box); 1/4 (\triangle); 1/8 (\diamondsuit) and 1/12 (\odot) in freestream wind speed (a) U_{∞} = 3.3 m sec⁻¹ and (b) U_{∞} = 6.6 m sec⁻¹. Also shown are the regressions by power-law fitting $\sigma_z = Ax^n$ (solid lines).



Fig. 6. Vertical dispersion coefficient σ_z plotted against friction length scale $L_f (= x^{1/2} \times \delta^{1/2} \times C_f^{1/4})$ in freestream wind speed $U_{\infty} = 3.3$ m sec⁻¹ (\blacklozenge) and $U_{\infty} = 6.6$ m sec⁻¹ (\blacklozenge). Also shown is the linear regression (solid line) for all the data points (the coefficient of determination by liner regression $R^2 = 0.933$). Size of roughness elements *h* is used as the characteristic length scale.

Supporting Information



Fig. S1 Measurement locations in streamwise direction.



Fig. S2. Vertical dimensionless profiles of pollutant concentrations $\overline{\psi}/\overline{\psi}_{canopy}$ plotted against wall-normal distance (z-h)/h at different streamwise locations x/h over arrays of street canyons of aspect ratios ARs = 1/2 (\Box); 1/4 (\triangle); 1/8 (\diamondsuit) and 1/12 (\circ) in freestream wind speed $U_{\infty} = 3.3$ m sec⁻¹. Filled symbols are the corresponding quantities in $U_{\infty} = 6.6$ m sec⁻¹.